

# Distribution of Depth and Clay-Silt to Sand Ratio of Land Subsidence in Coastal Semarang City by Resistivity Methods

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## Abstract

Semarang City is the capital of Central Java province, located in the northern coast of Java island, Indonesia which is geologically composed of alluvial deposits. The natural consolidation of young alluvial deposits has caused a land subsidence. On the other hand, load of buildings and constructions causes an increase in the pressure of land surface, and finally, leads to an increase in the rate of land subsidence. The drilling data indicated that not all layers of lithology are soft layers supporting the land subsidence. A research on the depth of soft lithology and its percentage of hard lithology has not been conducted ever before. On the other side, an analysis regarding this kind of research can be conducted based on resistivity method. Sounding system method with Schlumberger electrode configuration was selected for field data collection. The goals of this research were to determine the resistivity of lithology so that it can be used to interpretate distribution of depth and clay-silt to sand ratio of land subsidence area. The results showed that the area with the depth of the upper layer of clay-silt less than 3 m and clay-silt to sand ratio less than 4 were the high land subsidence area.

**Keywords:** Land subsidence, Semarang, Resistivity

## Introduction

Semarang is one of the the capital of Central Java province and nowadays suffering from extended land subsidence. Land subsidence, as a movement of a surface downwards relative to a datum such as sea level, is a major constraint to the development in many refions all over the world. As the effect of the land subsidence, some areas have been in below the sea level so it will be flooded with coastal flooding (rob) (Marfai and King, 2007). It can damage existing buildings, roads, bridges, industrial estates, and results in loss of homes, so it results in disturbing the activity of the people. The handling of this situation has been managed to be conducted by arranging the function of the area by paying attention to all behaviors of land subsidence, but it is not completely succeed yet.

The research on land subsidence rate in Semarang city has been carried out by experts and shows that there was a difference between one place and another. Based on the estimation from Levelling, Interferometric Synthetic Aperture Radar (InSAR), Microgravity and Global Positioning System (GPS) survey methods, land subsidence with rates of up to about 19 cm/year were observed during the period of 1999 up to 2011 (Gumilar *et al.*, 2013). That rate of land subsidence was affected by the process of alluvial consolidation and added by the exploitation of overflowing ground water and the building construction in some points (Abidin *et al.*, 2010; 2012). Tectonic activity is not related to the land subsidence so that the bedrock does not decrease (Wardhana *et al.*, 2014) The effect of land subsidence is predicted to cause the flooded area increases from 2,162.5 Ha (5.6%) to 3,896.3 Ha (10.1%) in the

2012 up to 2022 (Suhelmi, 2012). The rate of land subsidence in Semarang City is generally getting higher in the location with the thickness of soft lithology layer detected as the lithology with resistivity  $< \text{ohm}$  (Widada *et al.*, 2017<sup>a</sup>). Eventhough, it was not yet explained the other lithology in the layer system and it's distribution which is contributing in the land subsidence.

This research aimed to map the lithology layer system which resulting in the land subsidence by identifying thickness comparison of clay-silt and sand and also the cover layer thickness of the soft lithology layer and mapping the spread spatially according to the data of resistivity measurement.

**Materials and Methods**

Resistivity method is measurement technique of geophysics that uses resistivity or electricity characteristic of rocks to observe subsurface of earth condition (Widada *et al.*, 2017<sup>a</sup>). Measuerment was conducted in the land surface by injecting electricity current (in miliAmpere) along two current electrodes (C1

and C2) then gaining potential difference (in miliVolts) from two potential electrodes (P1 and P2). There is positive connection between rare electrode with the current penetration into the earth. The widest electrodes space could be get a depth currecnt penetration, hence the physical characteristics of deeper rock layer can be analyzed (Setyawan *et al.*, 2016). The location points of resistivity measurement can be seen in Figure 1.

In the field implementation using Schlumberger configuration (Dobrin, 1976). The distance of the electrode was arranged so that  $R1 = R4 = (b - \frac{1}{2} a)$  and  $R2 = R3 = (b + \frac{1}{2} a)$ , where  $a$  is the distance between the two potential electrode and  $b$  is the distance of the center point to the current electrode (Figure 2). In this state, the range of potential electrode P1 – P2 is started from 1/3 of range of the current electrode C1 – C2. Furthermore, the measurement was done by only moving the current electrode to a distance where the measurement result of the potential difference P1 – P2 is already small, then P1 – P1 is extended step by step.

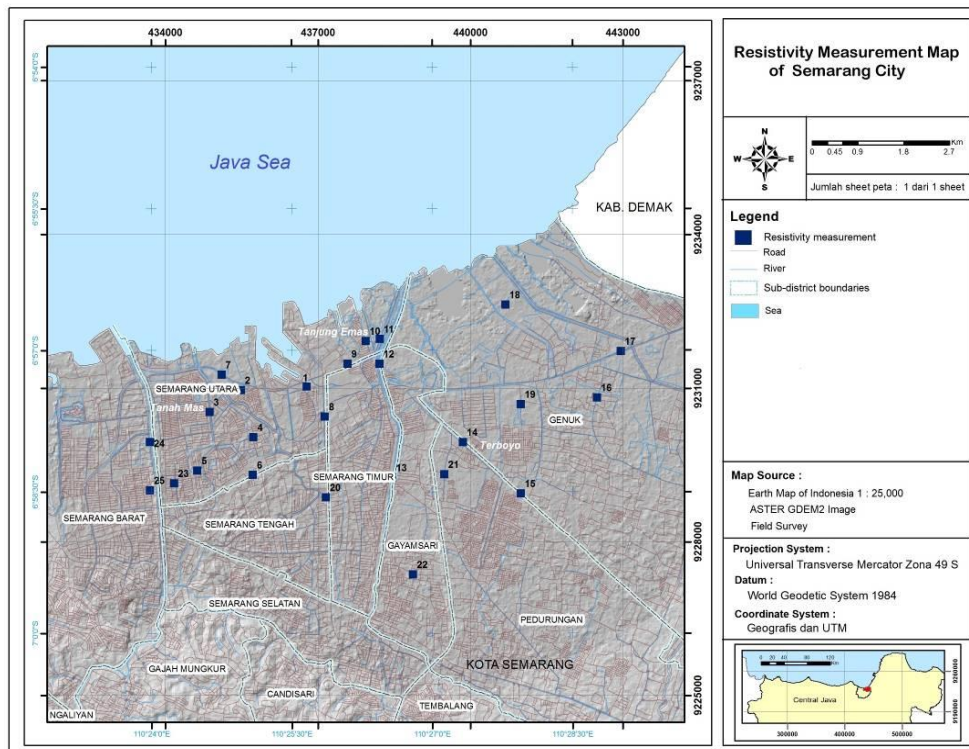
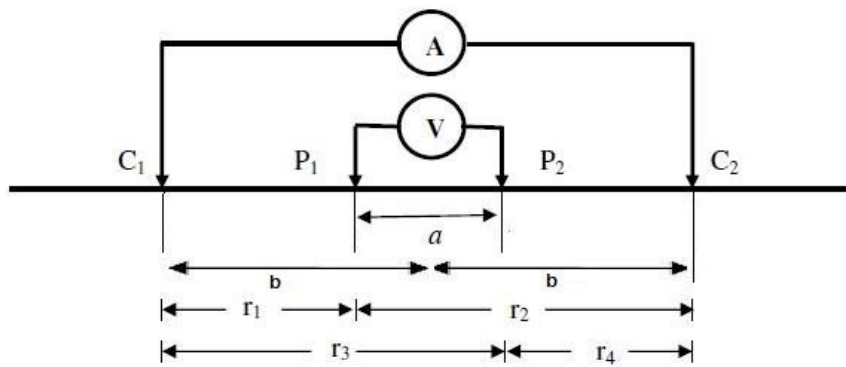


Figure 1. Resistivity measurement map



**Figure 2.** The scheme of electrode arrangement of Schlumberger configuration (Dobrin, 1976)

The field data obtained from the resistivity measurement were the distance of current electrode, distance of potential electrode, current ( $i$ ), and the potential difference ( $\Delta V$ ). The resistivity of the medium was calculated from the quotient of the  $\Delta V/i$  using a modified Ohm's law. Based on the data, the apparent resistivity was measured for every current injection, so the value of apparent resistivity was obtained for every electrode spread. The graph of the spacing electrode and the apparent resistivity then were used as the basic in determining the thickness and the true resistivity by using the Ip2Win software. The results obtained were the thickness and the resistivity of every lithology layer (Widada *et al.*, 2017<sup>a</sup>)

The resistivity for every lithology layer that was obtained in the data processing, was then interpreted the type of lithology by referring to the resistivity of lithology type from the references and correlated to the type of lithology based on the geotechnical drilling data. In the geotechnical drilling data, besides the description of the type of lithology, also listed the physical characteristics of lithology related to its ability to accept the burden, so it could be interpreted the behaviors of lithology related to the land subsidence. From the interpretation result in every point of resistivity measurement, a contour map was made to depict the depth of soft lithology (clay and silt) layer which causes the land subsidence in the research area. Comparison of clay-silt thickness to sand which has a greater bearing capacity was also analyzed to get an overview of its influence on land subsidence.

## RESULT AND DISCUSSION

Based on Semarang and Magelang Geological Map (Thanden *et al.*, 1996) Semarang has three main lithologies, namely, volcanic rock, sedimentary rock, which is marine in origin, and alluvial deposits. Volcanic rock consists of volcanics breccias, lava flow, tuff, sandstone, and clay stone. This area is located on the Southern part of Semarang. Sedimentary rock originated from marine consists of clay stone and dominated by sandstone in between. Alluvial sediment consists of beach deposits formed by clay and sand with a thickness of more than 80 m with the age of the Holocene period. It is on these deposits (Marine sediment and alluvial sediment) where the land subsidence is occurring. The layer of clay and silt will prone to have smaller resistivity than the sand, especially if it contains water (Yulianto *et al.*, 2016). Resistivity of clay and silt in Semarang alluvial deposits was about less than 3 ohm, while sand resistivity ranges from 5 – 20 ohm (Widada *et al.*, 2017<sup>b</sup>). Based on this criterion, the results of the depth of clay-silt layer and comparison the thickness of clay-silt to the thickness of sand (clay-silt to sand ratio) were presented in Table 1. In this table, each resistivity measurement point pairing with the data of land subsidence measurement (Wardhana *et al.*, 2014).

The map of land subsidence overlaid by the map of depth of top soft layer clay-silt is presented on Figure 3. Based on the figure, it can be seen that the rate of land subsidence at the location around Tanah Mas, Tanjung Emas Port and Terboyo are higher than another areas (brown and red

color). In these areas, depth of top soft layer are less than another areas (shaded vertical and horizontal line). This pattern indicates that the thinner layer of the soft layer covers, then the burden of the building on it rests directly on the soft layer. Large pressure on the soft layer enlarges compression and increases in land subsidence.

The location around Tanjung Emas Port and Tanah Mas in North Semarang Sub-district had the higher land subsidence than the other areas at 9.4 - 10.08 cm/year. In that location, it was found that the depth of soft layer was also less than another areas at 1.0 – 1.5 m in Tanah Mas and 0.9 – 1.0 m in around Tanjung Mas Port. In the point 23 of resistivity measurement located in Krobokan with the land subsidence of 10.69 cm/year also had the depth of soft layer about 4.0 m. The

location with a slightly higher rate of land subsidence was also found in around Terboyo, at 8.58 cm/year, and it actually had the depth of soft layer for 2.3 m. The spread of soft layer can be found in all research areas with the depth varied 0.9 m to 10.8 m.

The map of land subsidence overlaid by clay-silt ratio is presented on the Figure 4. Based on the figure, it can be seen that the rate of land subsidence in the location around Tanah Mas, Tanjung Emas Port and Terboyo are higher than another areas (brown and red color). Generally, clay-silt to sand ratio in this area are less than another areas (shaded vertical and horizontal line). This pattern indicates the sand layer plays a role in minimizing land subsidence. A high land subsidence is actually happening in the thick soft layer. It happened because that

**Table 1.** The Depth of soft lithology layer, clay-silt to sand ratio based on the resistivity data and the rate of land subsidence in Semarang City

Point No.	East	North	Depth of Soft Lithology layer ( m )	Clay-Silt To Sand Ration	Rate of land subsidence (cm/year)
1	433697	9229943	8.0	1.273	5.77
2	435485	9230954	1.5	1.698	3.50
3	434868	9230527	3.0	4.241	5.28
4	435731	9230038	5.7	2.106	8.51
5	434629	9229383	1.0	3.137	10.69
6	435719	9229303	9.7	0.307	0.47
7	435105	9231261	1.0	1.873	9.90
8	437136	9230440	1.0	1.618	3.50
9	437583	9231474	1.0	1.134	10.45
10	437942	9231918	4.0	17.228	3.50
11	438217	9231957	0.9	0.414	9.80
12	438215	9231470	1.0	0.712	10.45
13	438464	9229248	6.2	1.499	7.00
14	439857	9229941	6.6	0.357	7.00
15	440994	9228939	10.9	2.447	5.09
16	442496	9230815	11.8	1.308	5.09
17	442960	9231722	2.2	0.347	8.81
18	440693	9232628	2.2	12.871	8.58
19	440992	9230686	3.3	0.541	5.76
20	437161	9228864	4.0	0.099	4.02
21	439486	9229312	4.5	1.729	7.00
22	438870	9227356	5.2	0.608	5.09
23	434170	9229137	4.0	0.445	10.69
24	433697	9229943	3.0	0.570	9.90
25	433697	9229000	8.0	1.273	0.50

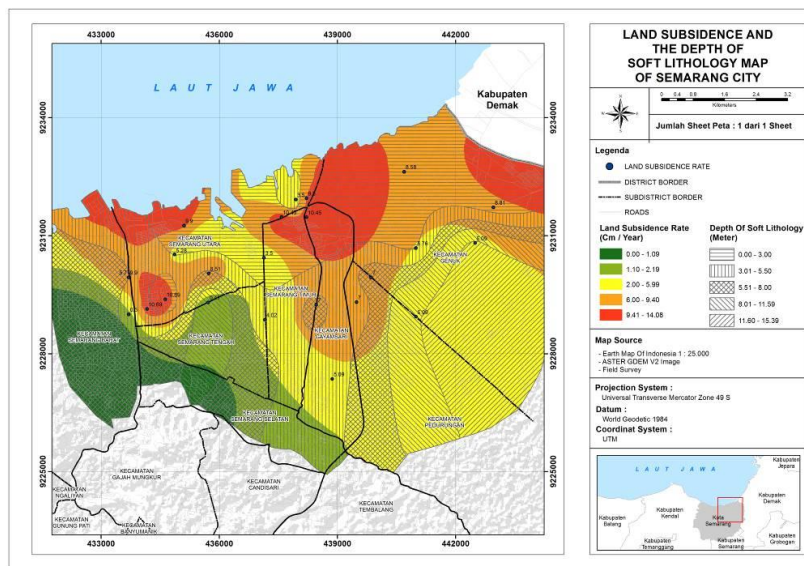


Figure 3. Map of the land subsidence and the depth of soft lithology in Semarang City

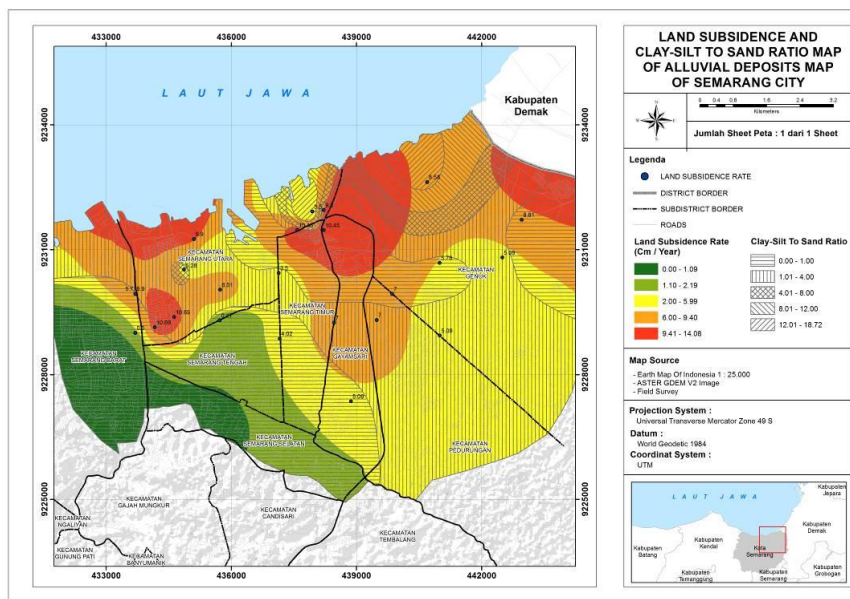


Figure 4. Map of the land subsidence and the clay-silt to sand ratio in Semarang City

soft layer is compressed when given the load from the above layer or from several buildings and infrastructures above the ground (Gumilar et al., 2013; Abidin et al., 2012). The consolidation process in the alluvial deposit is natural and surely happen (Reddish and Whittaker, 1989). Figure 4, also shows that there are an anomalies from the general pattern, namely in the eastern area of Tambaklorok to Terboyo, high clay-silt to sand ratio coincides with high land subsidence

area. This is possible because in that location a thick soft layer is close to the surface, so the presence of sand layers is less able to play a role in reducing land subsidence

### CONCLUSION

The rate of land subsidence in Semarang City is generally greater in areas with a small depth of soft layer. The small depth of soft layer in Tanah Mas, Tanjung

Emas Port and the area around Terboyo were about 0.9 – 2.3 m, and the rate of land subsidence was about 9.4 - 10.08 cm/year, higher than other areas. On the other hand, clay-to silt ratio in this area was 1,0 – 1,5 less than the another area. The anomaly shown in Tambaklorok-Terboyo area, areas with large land subsidies coincide with large silt to sand ratio. In this area thick soft layer is close to the surface, so the presence of sand layers is unable to play a role in reducing land subsidence.

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